

Contemporary strain rates in Fennoscandia from BIFROST GPS

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Baseline Inferences From Fennoscandian Rebound Observations, Sealevel and Tectonics

- *Hans-Georg Scherneck, Chalmers*
- *Rüdiger Haas, Chalmers*
- *Martin Lidberg, National Land Survey of Sweden*
- *Jan Johansson, Chalmers and SP Swedish National Research and Testing Institute*
- *Markku Poutanen, Hannu Koivula, Finnish Geodetic Institute*
- *Oddgeir Kristiansen, Norwegian Mapping Agency*
- *Glenn A. Milne, Univ. Durham, now Univ. Ottawa.*
- *James L. Davis, Harvard-Smithsonian Center for Astrophysics, Cambridge, Mass.*



Continuous GPS, daily solutions since Sep. 1993 using (16 ...) 64 stations

Done or in progress:

Glacial Isostatic Adjustment (GIA)

Model inversions for earth rheology and ice history
(examining 3D displacement rates)

Observed absolute sea level change 1.2 ± 0.2 mm/yr ----->



New:

Examine strain rates ("tectonics")

feasible, owing to the increasing consistency of observations:

- time series length, stable observing conditions
- IGS improvements, maturing systems
- (scale parameter rates may be a problem, translations and rotations not)



Would like to have a tool for the interpretation of the deformation field, especially the GIA residual

- as general as possible
(as an alternative to a specific model like GIA)
- emphasising the lithosphere
- emphasising (the not so well examined) horizontal motion
- enhancing anomalous regions (strain accumulation)

A: Strain rates

A-a: Using collocation (statistics)

A-b: Using a mechanical model

B: Something else (?)



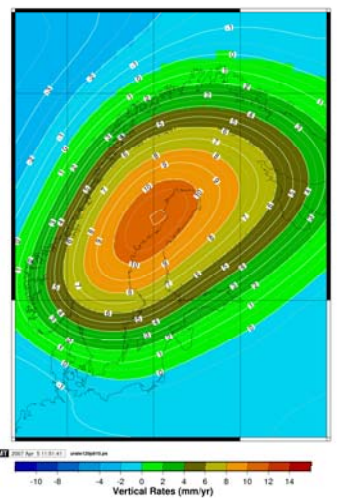
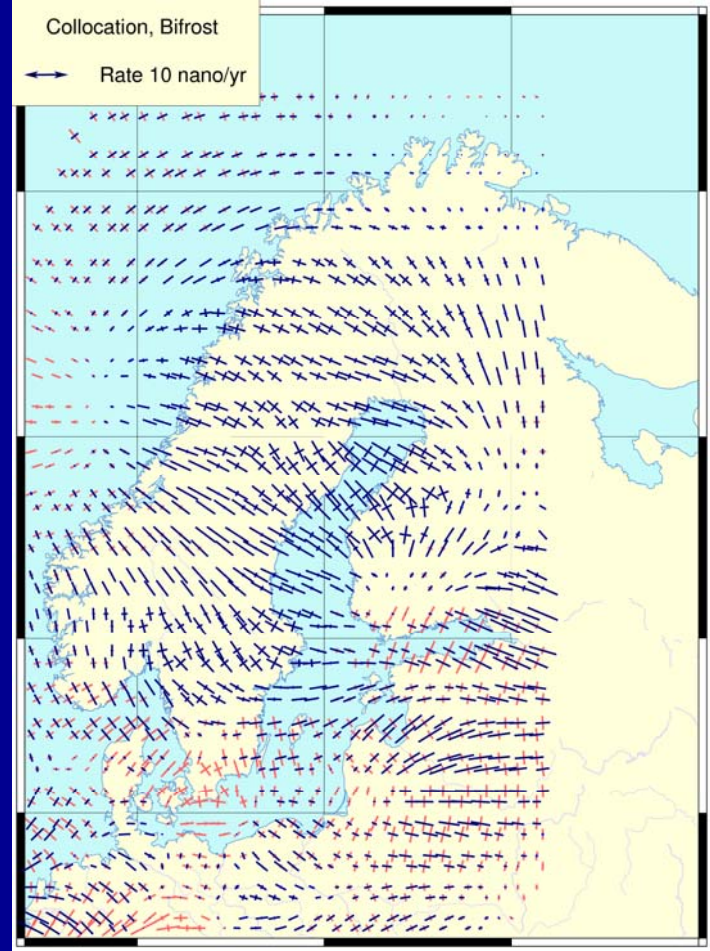
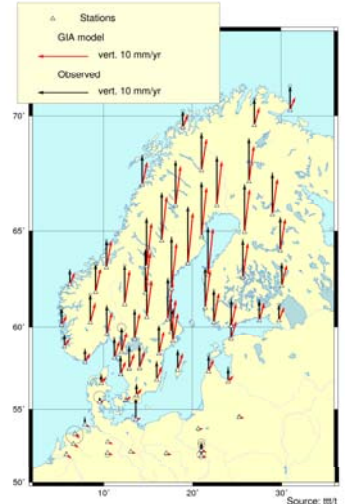
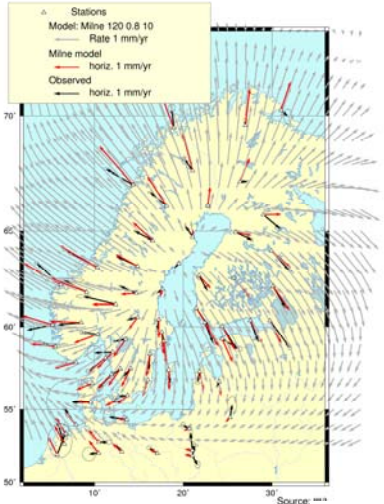
General: Leclerc (right)



something else



Collocation



Thick-plate concept

Motivation:

Collocation treats strain as a stochastic parameter

An elastic plate would reduce freedom to a physically consistent field

A thin plate would not make use of the vertical velocity component as all stations are on a surface

After all the lithosphere *is* thick, 200 km versus 2000 km diameter of rebound area

Flexure is important if vertical velocity = 10 mm/yr at wavelengths < 1000 km



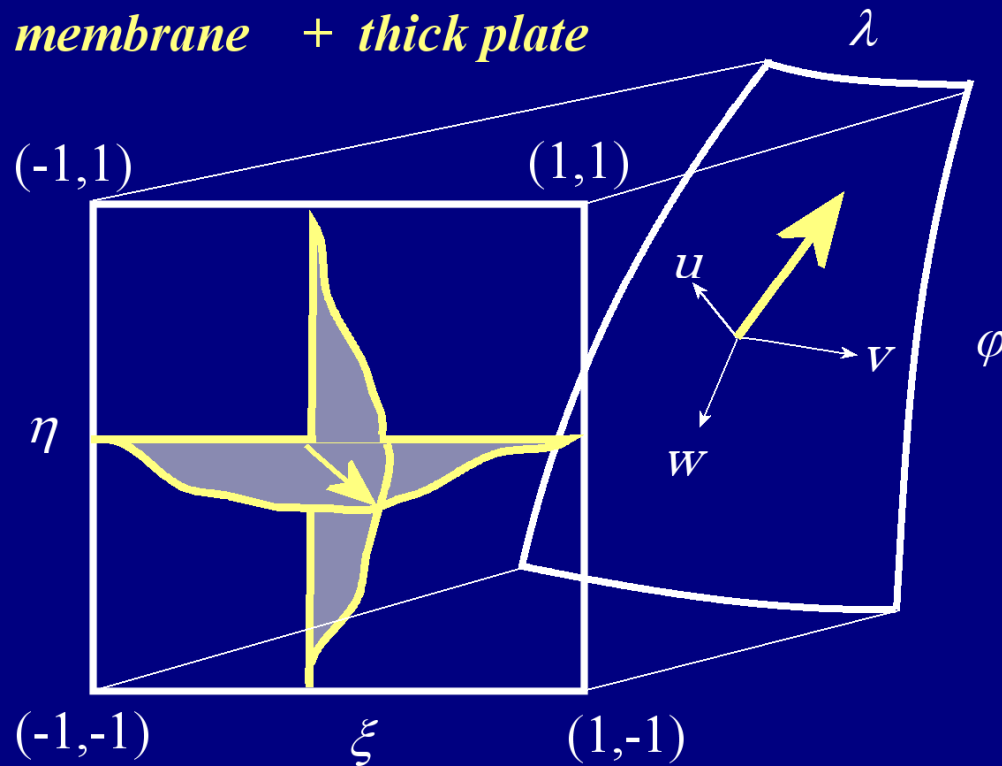
BIFROST project

"Finite Element" [sic!]: one element, high polynomial order, free-slip boundary conditions

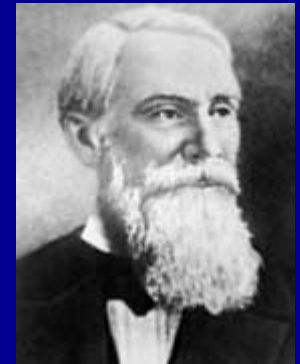
$$(u, v, w) = \sum_{i,j} (U_{ij}, V_{ij}, W_{ij}) T_i(\xi) T_j(\eta)$$

$$\epsilon_{\phi\phi} = \frac{1}{R} \frac{\partial w}{\partial \phi} + \frac{u}{R} - \frac{\zeta}{R^2} \left[\frac{\partial^2 u}{\partial \phi^2} - \frac{\nu}{\sin^2 \phi} \frac{\partial^2 u}{\partial \lambda^2} \right] \text{ etc.}$$

membrane + thick plate



$T =$
Chebychev
Polynomials



$$\epsilon_{\phi\phi} = \frac{1}{R} \left(\frac{\partial v}{\partial \phi} + u \right) \quad (2)$$

$$\epsilon_{\lambda\lambda} = \frac{1}{R \sin \phi} \left(\frac{\partial w}{\partial \lambda} + \sin(\phi)u + \cos(\phi)v \right) \quad (3)$$

$$\epsilon_{\phi\lambda} = \frac{1}{2R \sin \phi} \left(\frac{\partial v}{\partial \lambda} + \sin \phi \frac{\partial w}{\partial \phi} - 2 \cos(\phi)w \right) \quad (4)$$

$$\omega_z = \frac{1}{2R} \left(\frac{\partial w}{\partial \phi} - \frac{1}{\sin \phi} \frac{\partial v}{\partial \lambda} + \cot(\phi)w \right) \quad (5)$$



$$\epsilon_{\phi\phi}^{(\text{TP})} = -\frac{z}{R^2} \left(\frac{\partial^2 u}{\partial \phi^2} - \nu \frac{1}{\sin^2 \phi} \frac{\partial^2 u}{\partial \lambda^2} \right) \quad (8)$$

$$\epsilon_{\lambda\lambda}^{(\text{TP})} = -\frac{z}{R^2} \left(\frac{1}{\sin^2 \phi} \frac{\partial^2 u}{\partial \lambda^2} - \nu \frac{\partial^2 u}{\partial \phi^2} \right) \quad (9)$$

$$\epsilon_{\phi\lambda}^{(\text{TP})} = -\frac{z}{R^2} \frac{\partial^2 u}{\sin \phi \partial \phi \partial \lambda} \quad (10)$$



Cost function

$$P = \frac{1}{3N_O} \sum_{Obs} (\mathbf{u} - \bar{\mathbf{u}})^\top \Sigma^{-2} (\mathbf{u} - \bar{\mathbf{u}}) \\ + \frac{\Omega}{N_E} \sum_{E\text{-points}} \left\{ \int_{-D/2}^{D/2} \boldsymbol{\sigma} : \boldsymbol{\epsilon} dz \right. \\ \left. + g\rho_{eff}u^2 \right\} \\ = \frac{1}{3N_O} \chi^2 + \frac{\Omega}{N_E} (E_b + E_p + E_g)$$

Ω small: observation dominated

Ω big: strain energy dominated

Today: only observation dominated!

observation misfit

plane stress and flexure work

gravity work



Uncertainty analysis of inferred strain rates

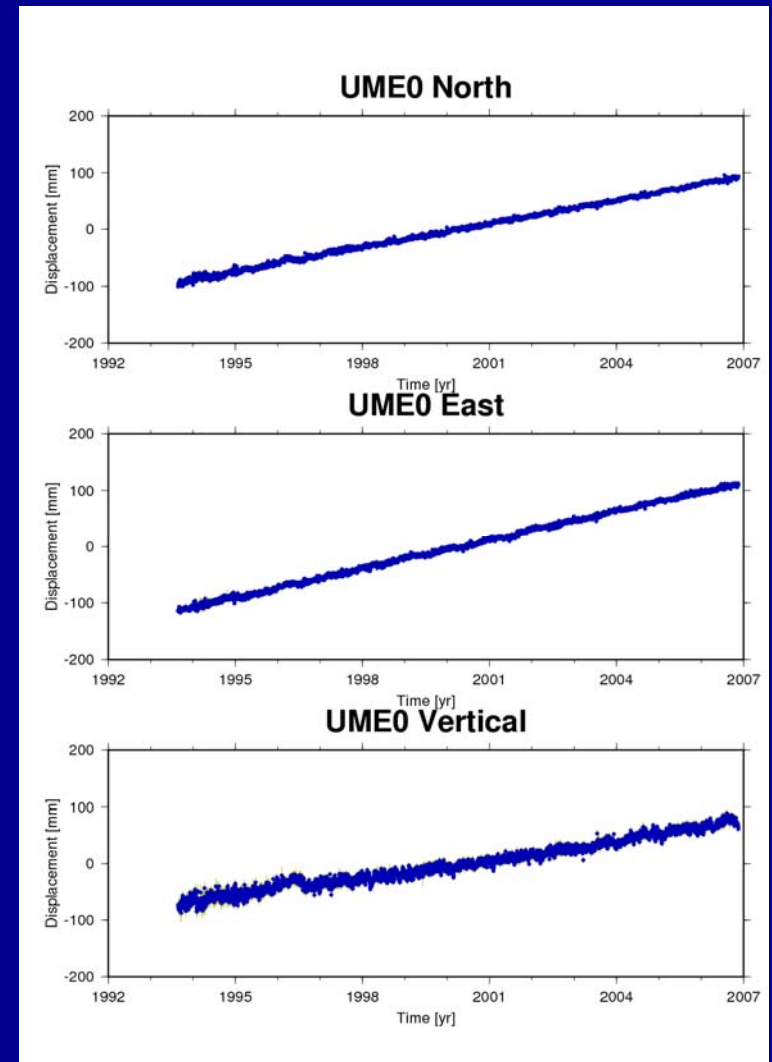
- Vary each Chebychev coefficient (say: j)
such that the χ^2 of the observation misfit doubles
(for 1 degree of freedom, this implies 85% confidence – N.B. we vary one coefficient at a time)
- Compute RMS:
Root-mean of
 $\{\text{deviation}_j \text{ of strain field}\}^2 / \text{number of coefficients of all } j$

A typical normalised χ^2 of fit of observations is 10



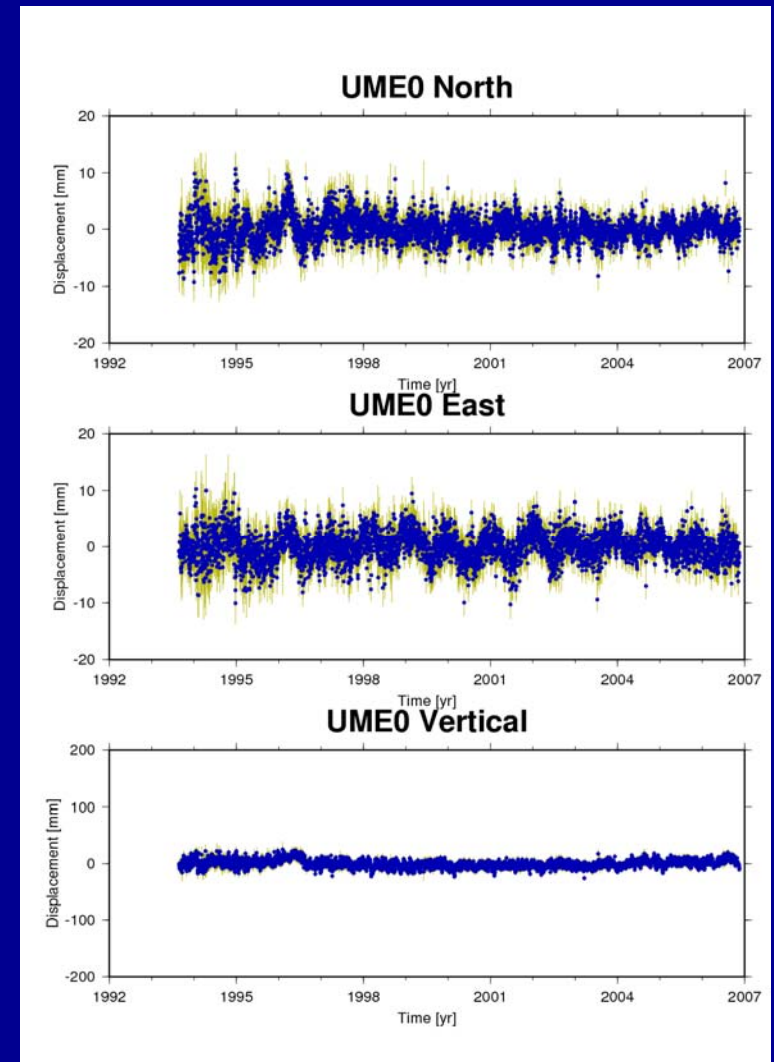
Observations

- GPS 1996-2007
 - daily solutions, 84 stations
 - GAMIT
 - ITRF2005
- + adjustment for vertical rates by local stabilisation
(Lidberg, PhD-thesis May 2007)



Observations

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----- BIFROST – GPS -----

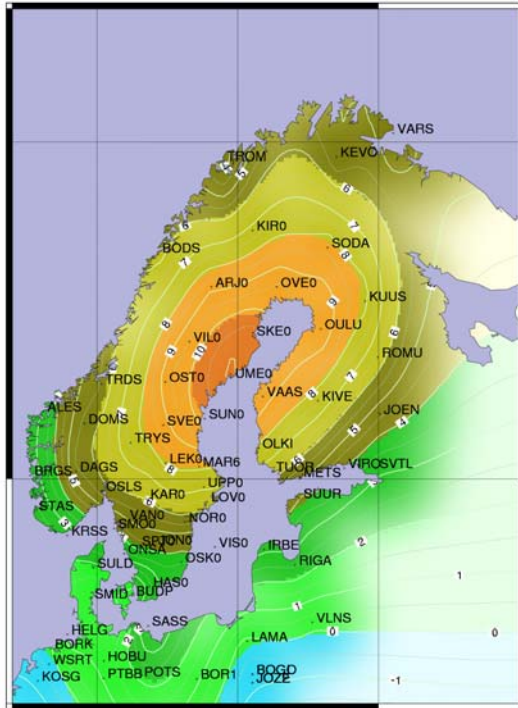
interpolated (raw)

Observed vertical

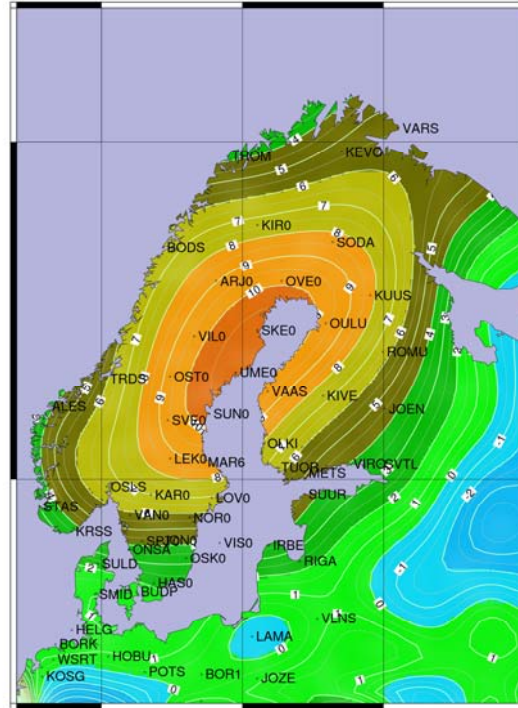
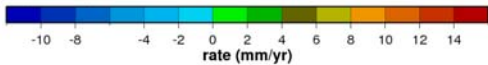
thick-plate model

Polyfit vertical

GIA model

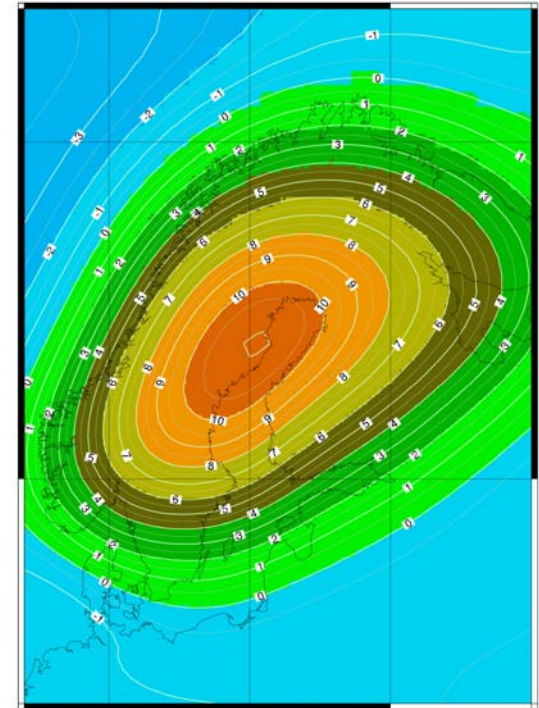
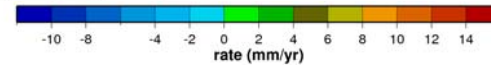


GMT 2007 Apr 11 19:51:14 urate-obs.ps



GMT 2008 Oct 3 15:43:56 ttpstf/urate.ps

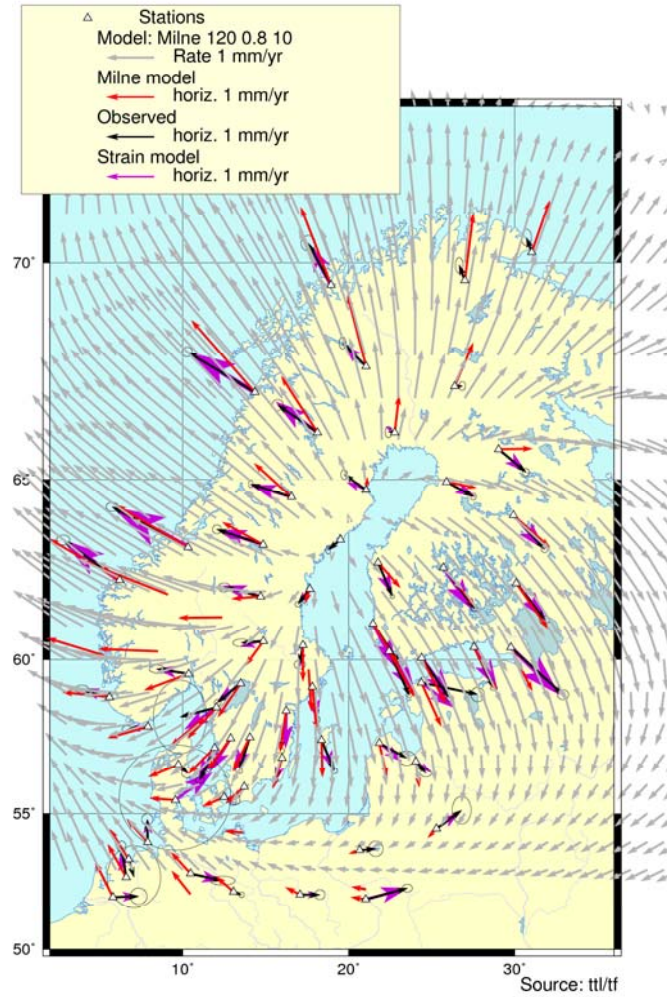
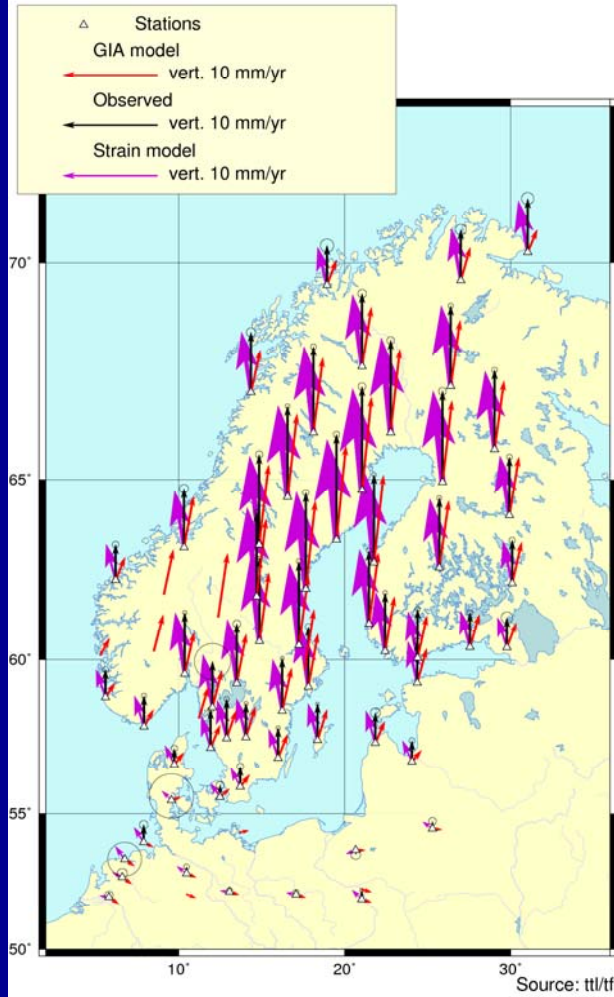
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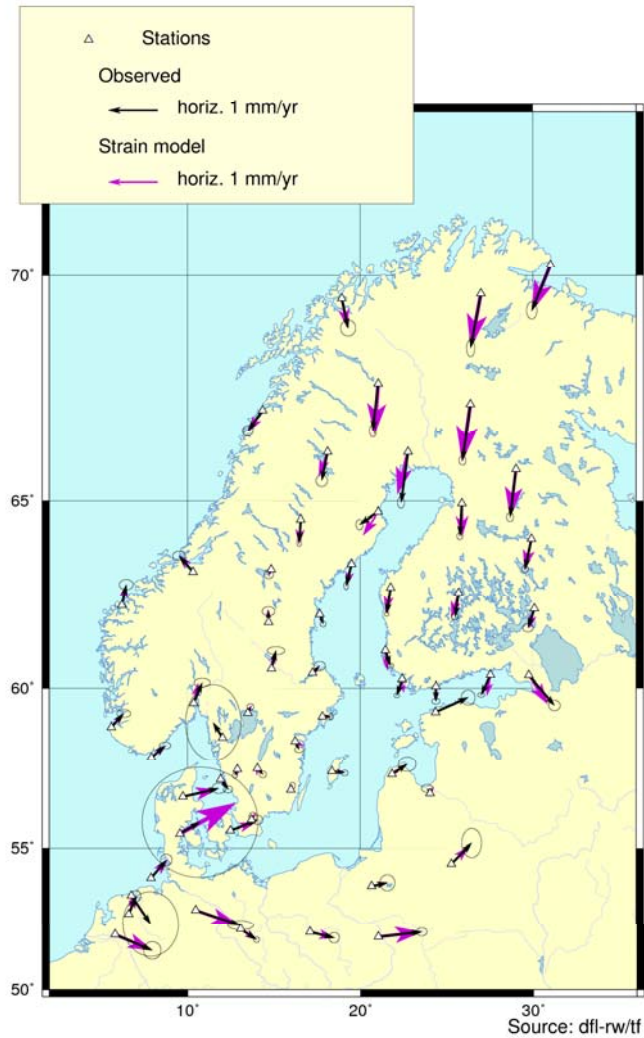
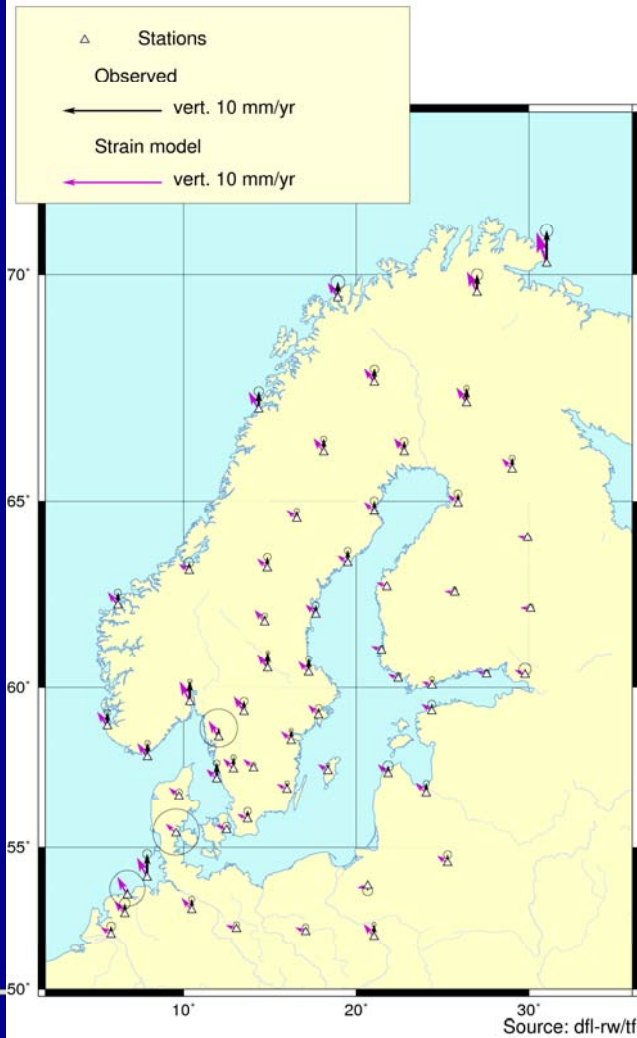
GMT 2007 Apr 5 11:51:41 urate120p010.ps



BIFROST project



"Observed" = GPS Observations – GIA model



Results of the strain analysis

Data sets:

- The GIA model on its own grid
- The thick plate model
 - GIA model
 - GPS
 - GPS-GIA residual

Products

- Areal strain
- Shear strain
- Curl
- Uncertainties for each (3)
- Strain energy: Plane and bending, areal and shear (4)
- Observation misfit

=> $33 + 3 = 36$ maps to show!



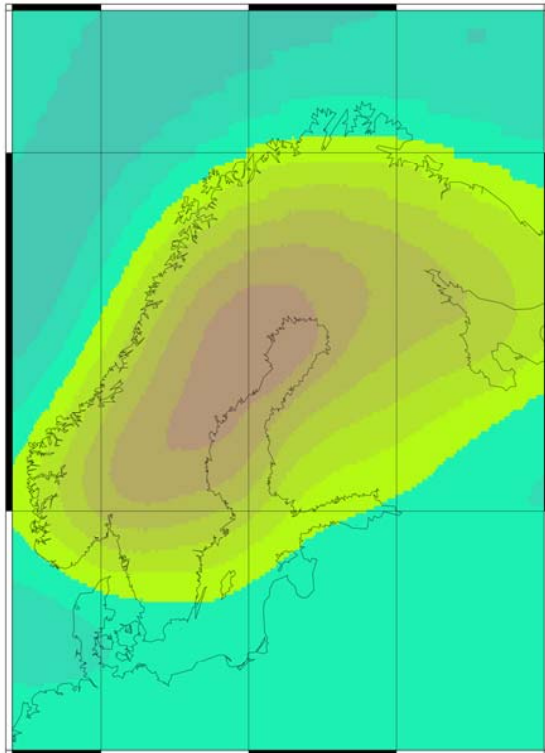
First showing...

- GIA Model on a regular grid $1/12 \times 1/12$
- Thick-plate solution to GIA model
 - Represented at the BIFROST sites
 - Using the observation uncertainties
- Strain and curl Uncertainty

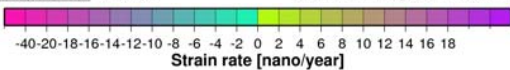


GIA model grid | --GIA model on BIFROST netw.-- |

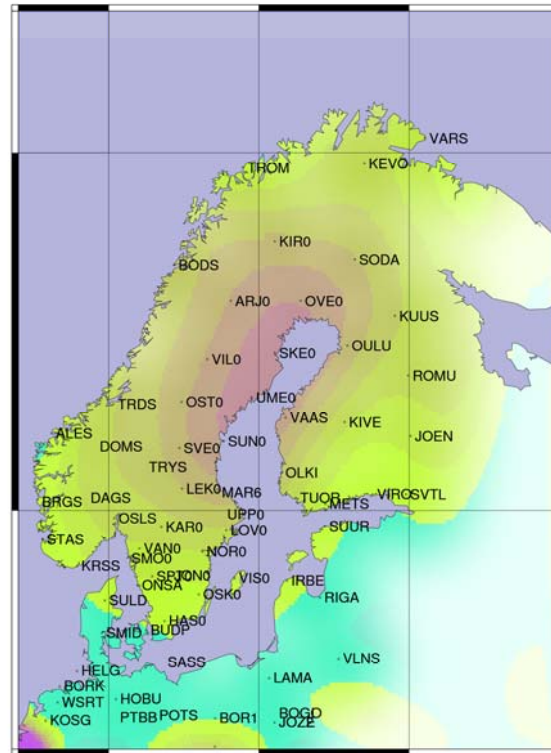
Areal Strain



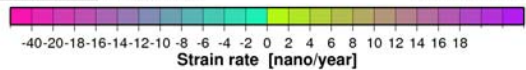
Source: Milne grids, 120 0.8 1



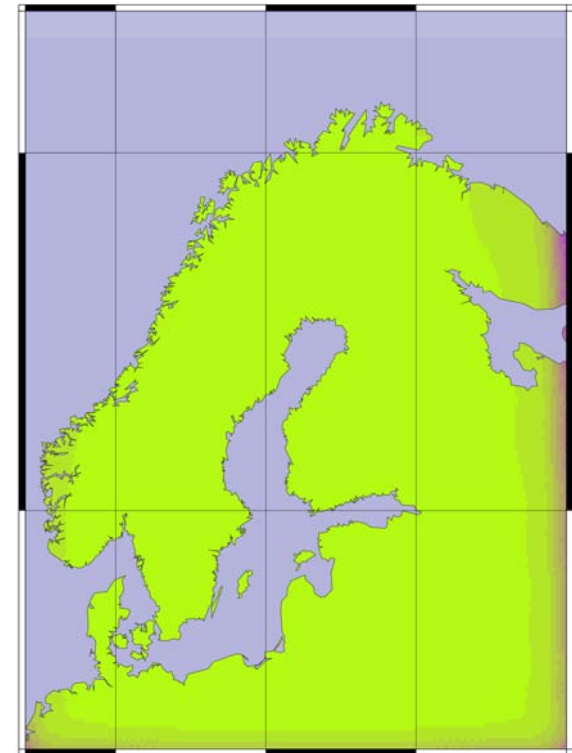
Areal Strain



Source: siml/TF



Areal strain



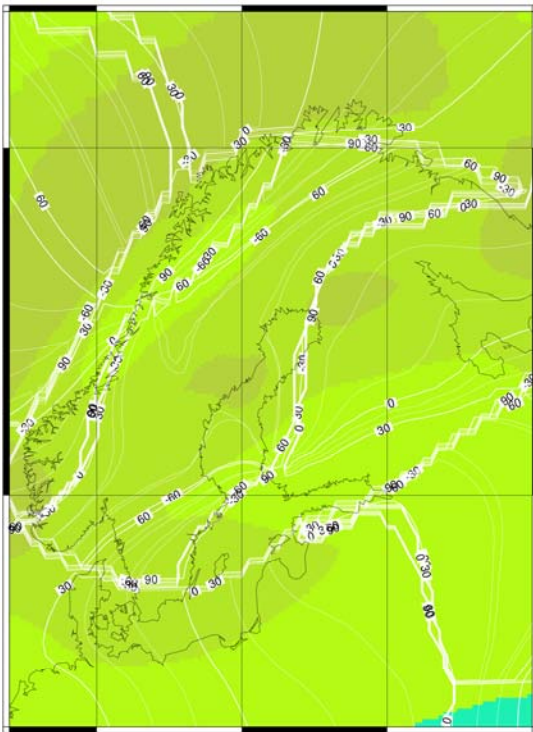
Source: siml/TF



GIA model grid

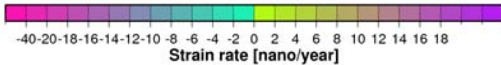
|--GIA model on BIFROST netw.--|

Shear Strain

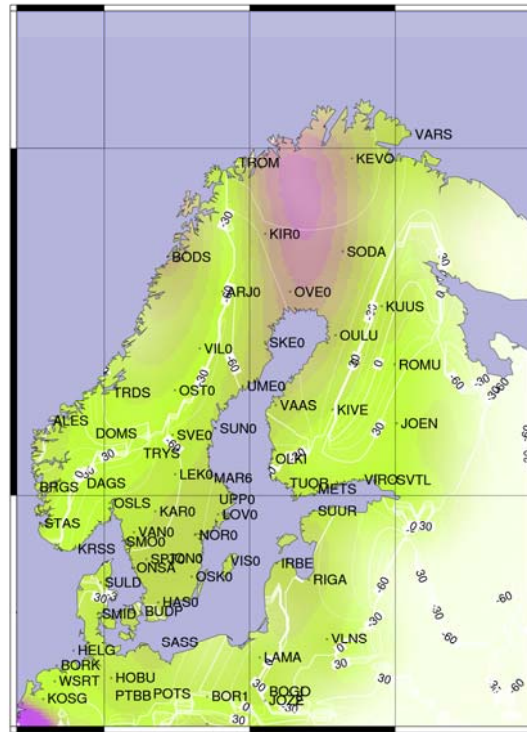


GMT 2008 Apr 9 13:19:09 shearstrain.ps

Source: Milne grids, 120 0.8 1

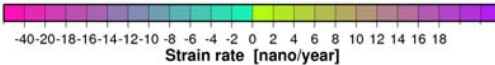


Shear strain

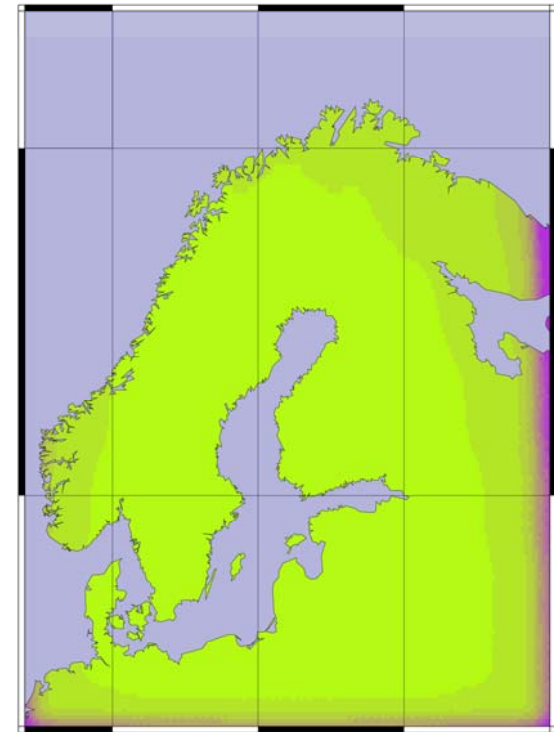


GMT 2008 Oct 2 15:26:30 simIPSTF/shearstrainr.ps

Source: sim/1f

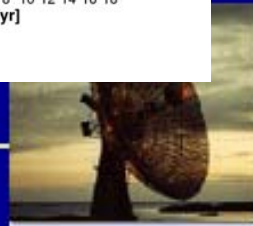
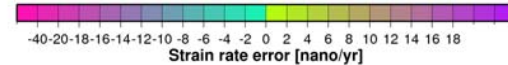


Shear strain



GMT 2008 Oct 2 15:26:31 simIPSTF/shearstrainr.ps

Source: sim/1f



BIFROST project

Curl = Rotation. Theoretically, on a radially symmetric earth loaded by normal forces there is no curl.

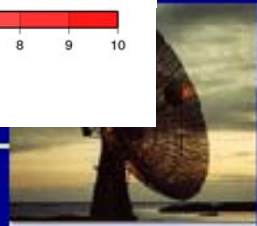
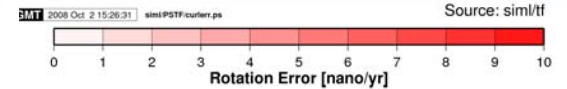
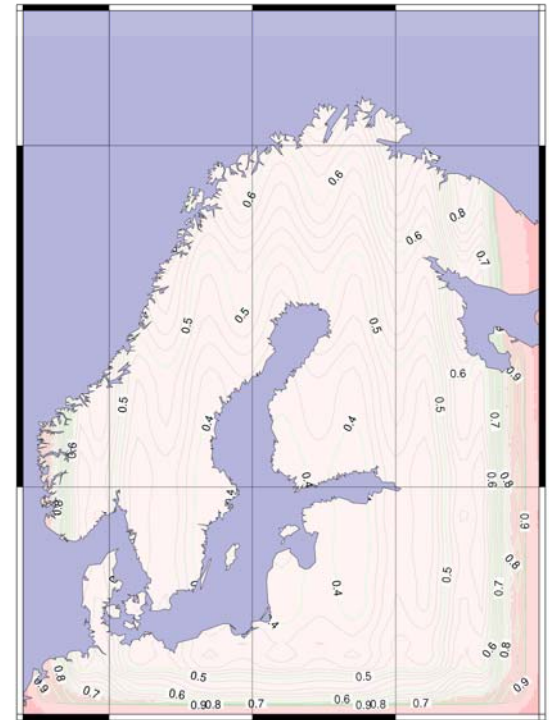
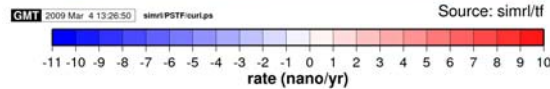
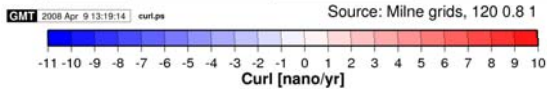
GIA model grid

----- GIA model on BIFROST netw. -----

Curl

Rotation

Rotation



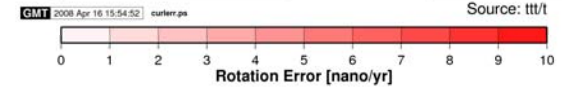
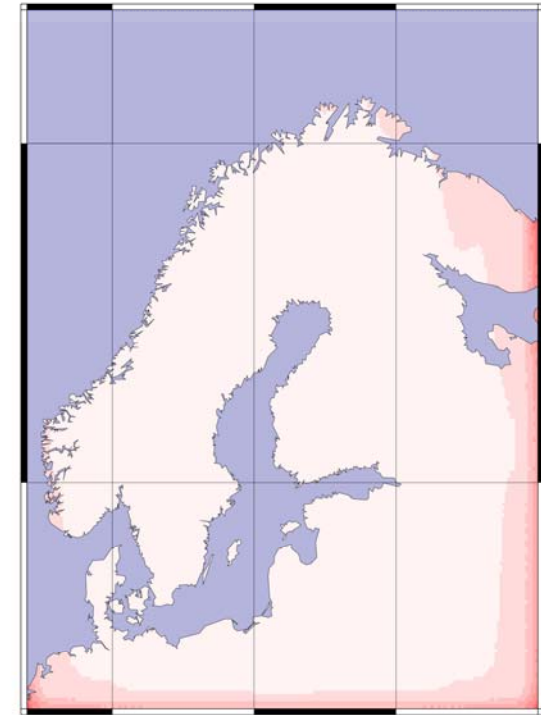
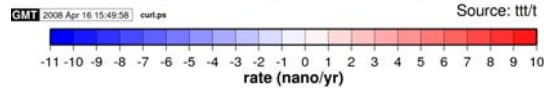
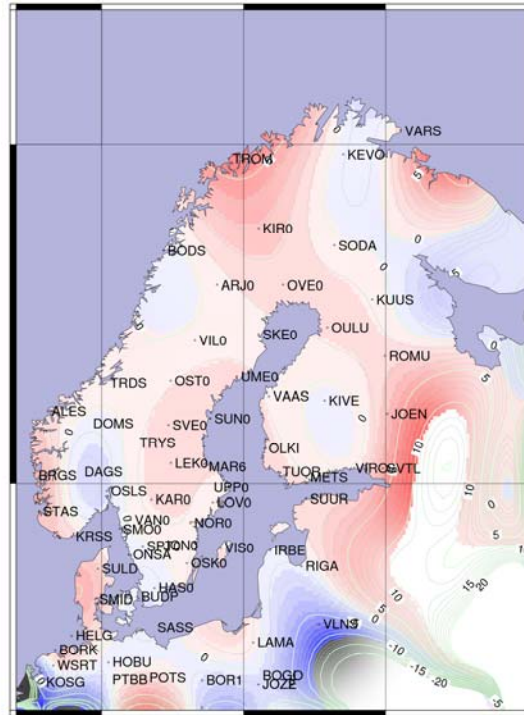
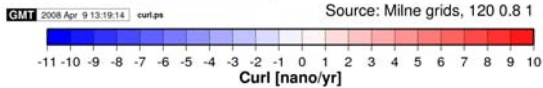
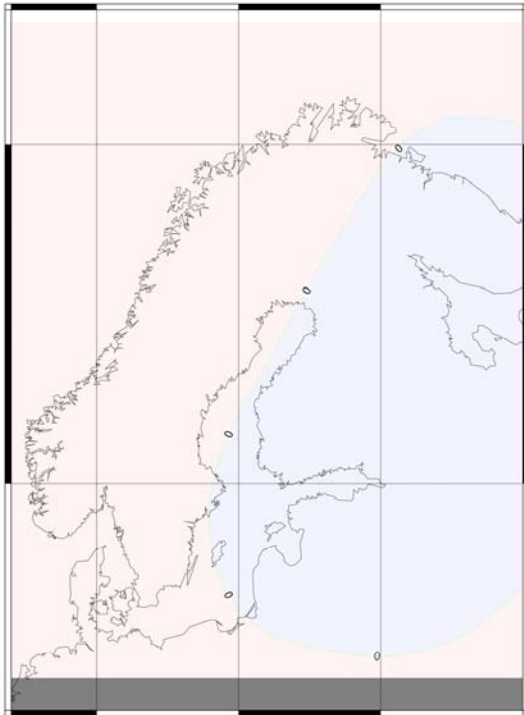
GIA model grid

----- BIFROST GPS -----

Curl

Rotation

Rotation



Nota bene

The curl does not vanish when the thick plate is fit to curl-free model data!

Similarly, shear strain is amplified

The curl would be one tectonic parameter, so there is no motivation to force it to zero or to issue an extra penalty.

The anomalies occur primarily where the station network is sparse

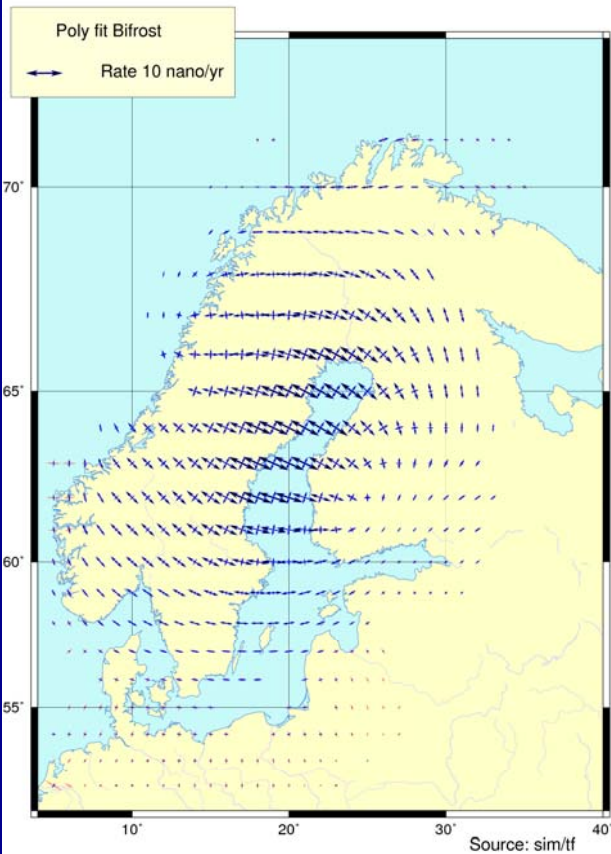


Input: GIA model data

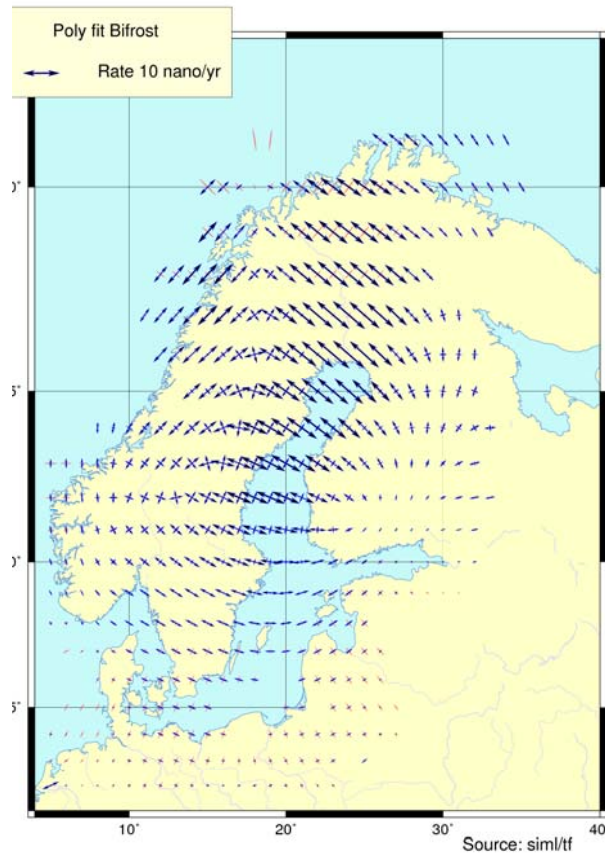
Thin plate

Thick plate

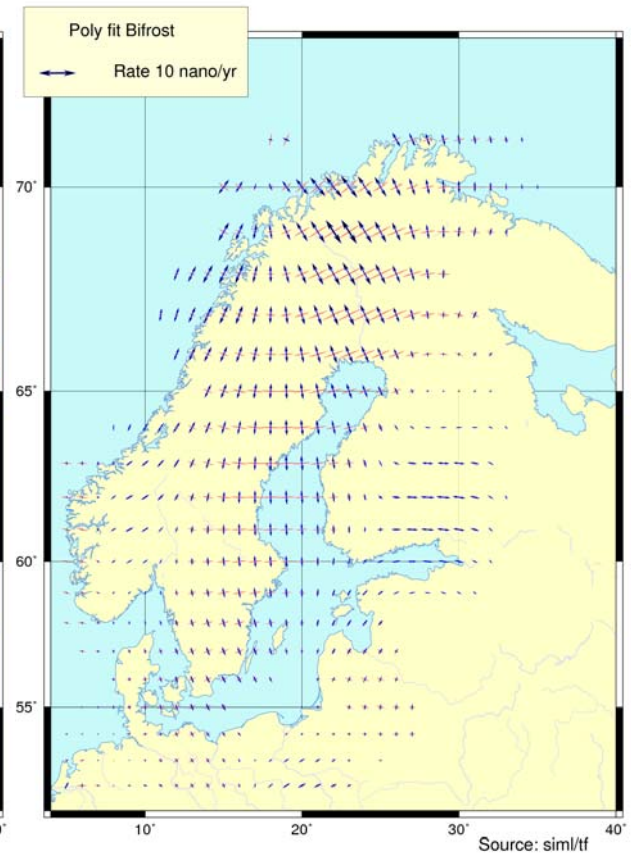
Surface strain rates



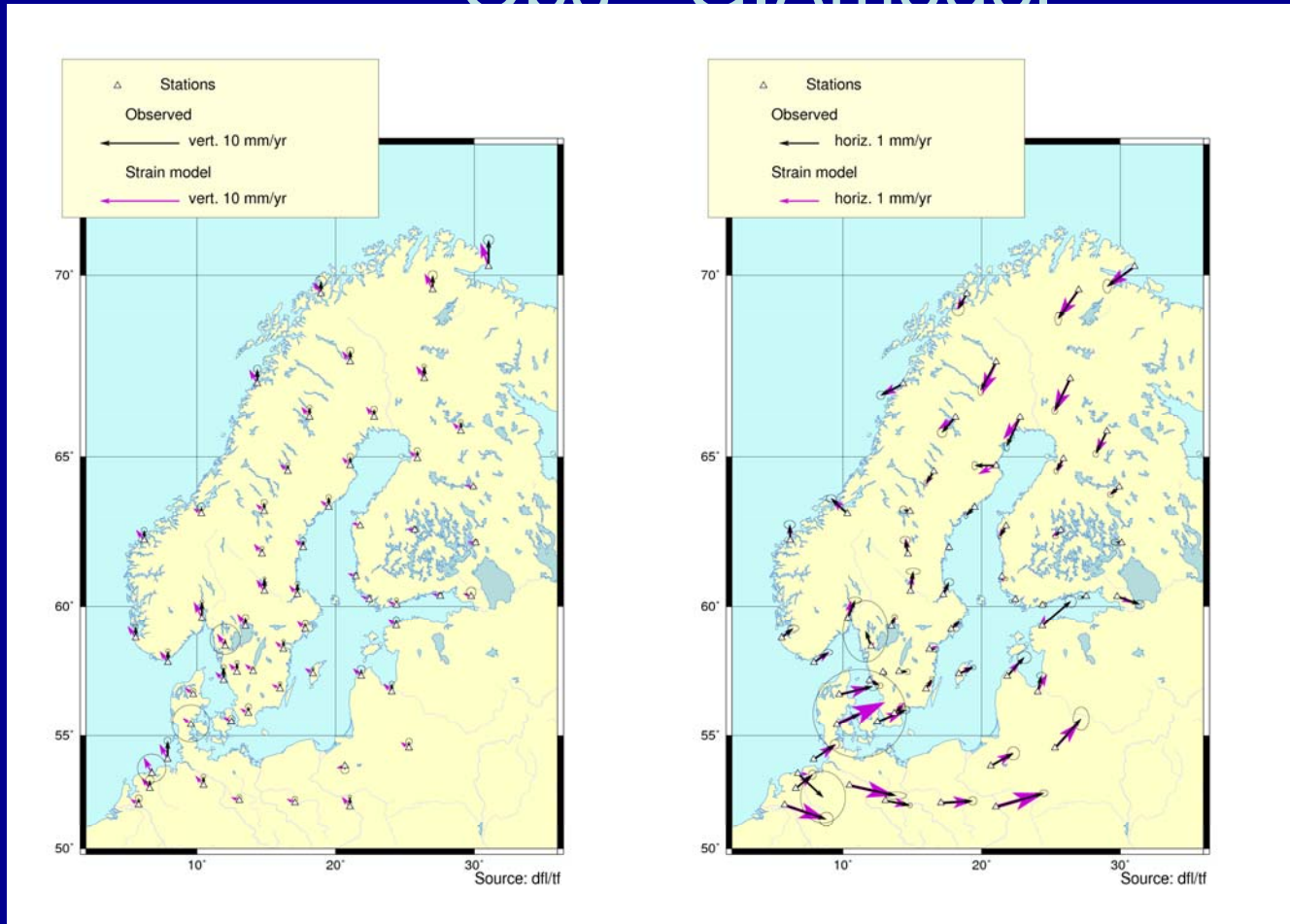
Surface strain rates



In-plane strain rates



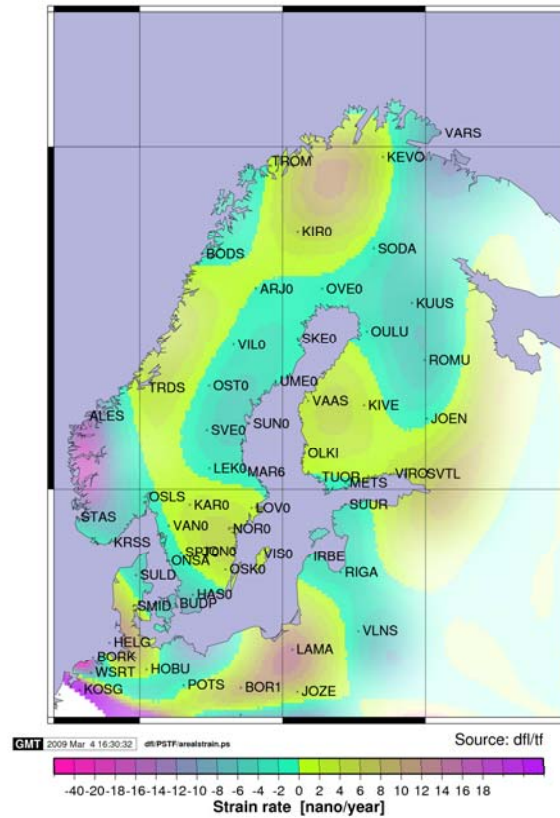
Next: Doing the analysis on the residual Obs - GIAmodel



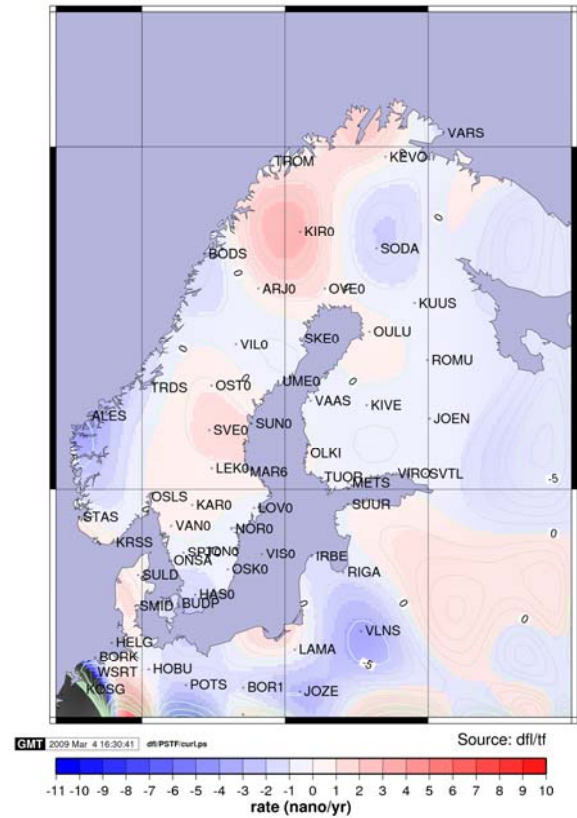
Surface strain rates



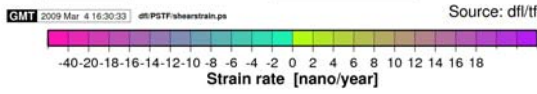
Areal Strain



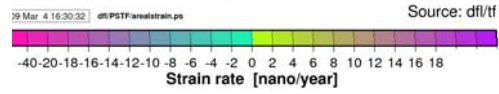
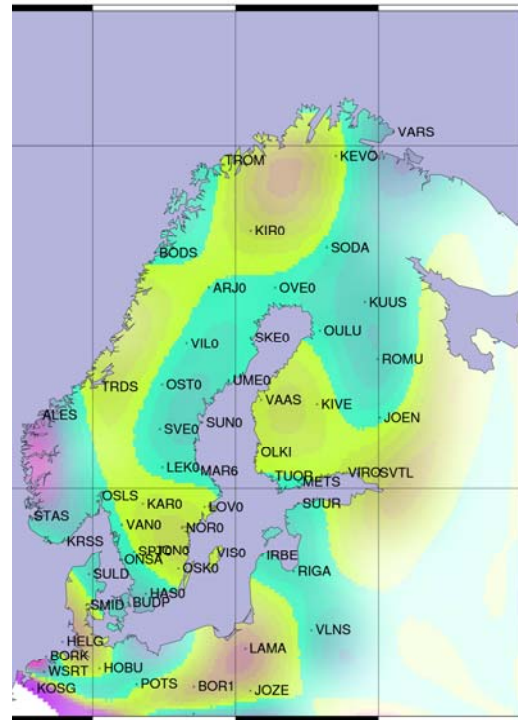
Rotation



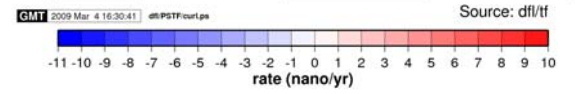
Shear strain



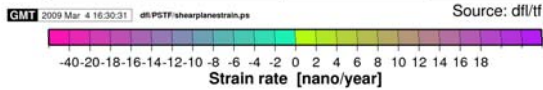
Areal Strain



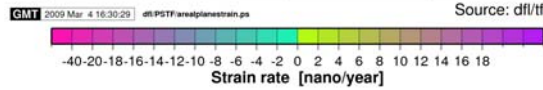
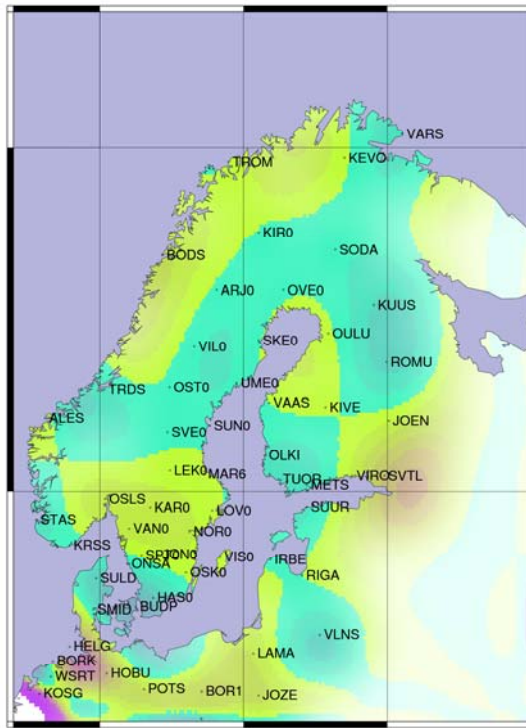
Rotation



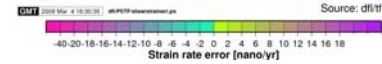
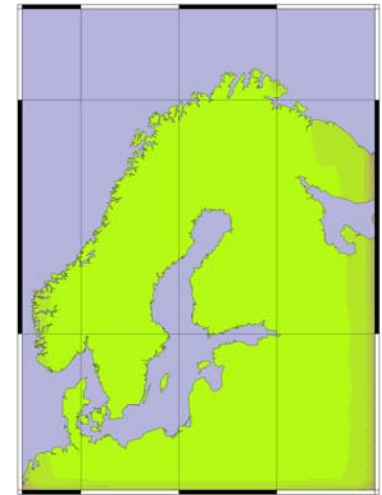
Shear strain in-plane



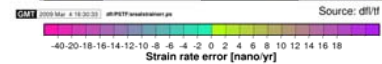
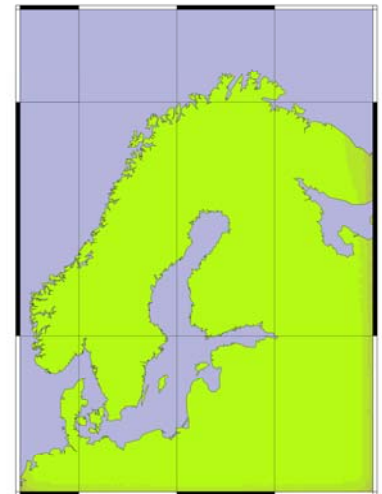
Areal strain in-plane



Shear strain



Areal strain

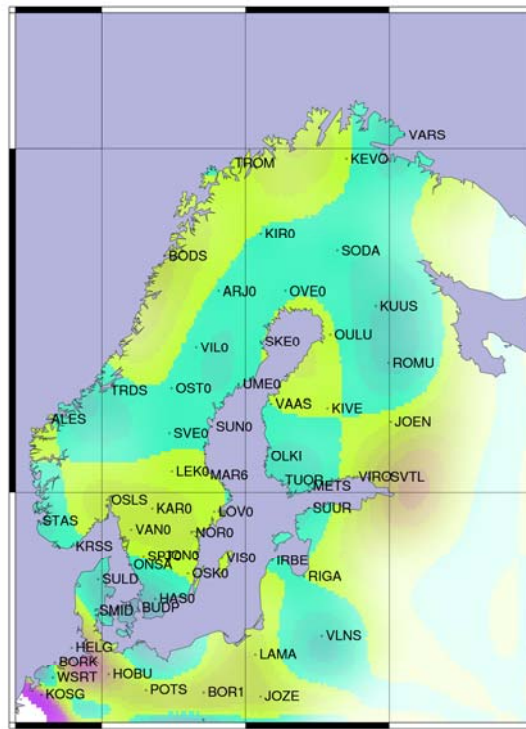


Shear strain in-plane



GMT 2009 Mar 4 16:30:31 d:\PSTF\shearplanestrain.ps Source: dfl/lf
 -40 -20 -18 -16 -14 -12 -10 -8 -6 -4 -2 0 2 4 6 8 10 12 14 16 18
 Strain rate [nano/year]

Areal strain in-plane



GMT 2009 Mar 4 16:30:29 d:\PSTF\arealplanestrain.ps Source: dfl/lf
 -40 -20 -18 -16 -14 -12 -10 -8 -6 -4 -2 0 2 4 6 8 10 12 14 16 18
 Strain rate [nano/year]

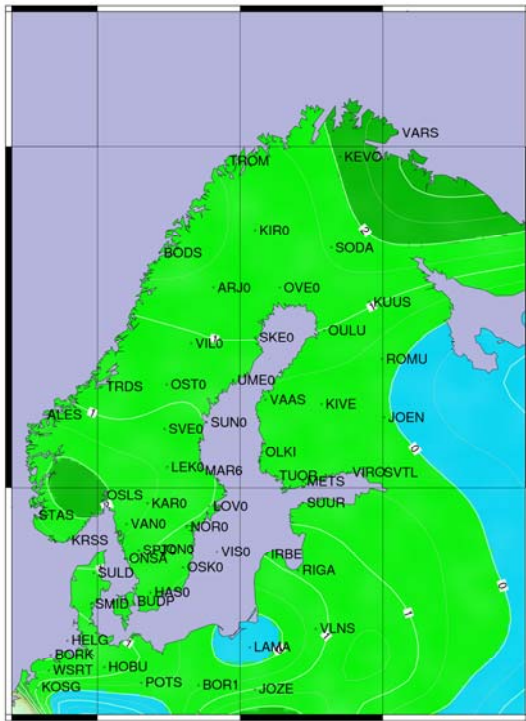
In-plane strain rates



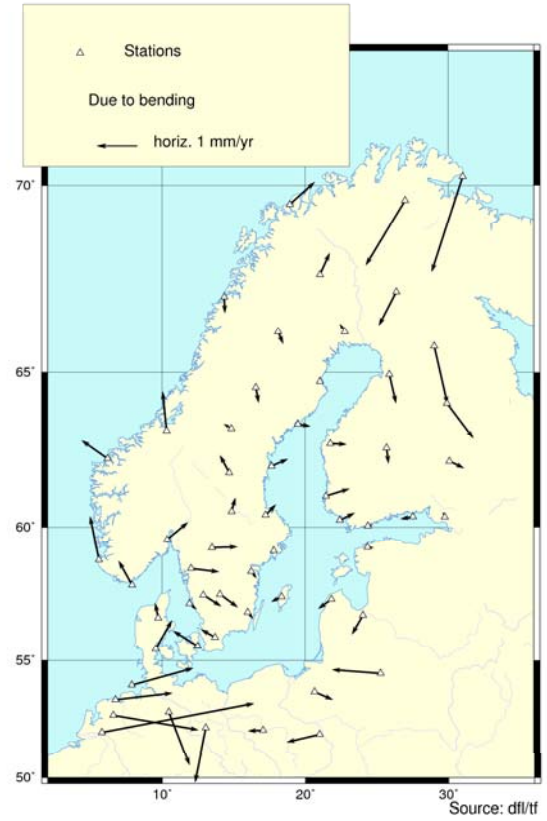
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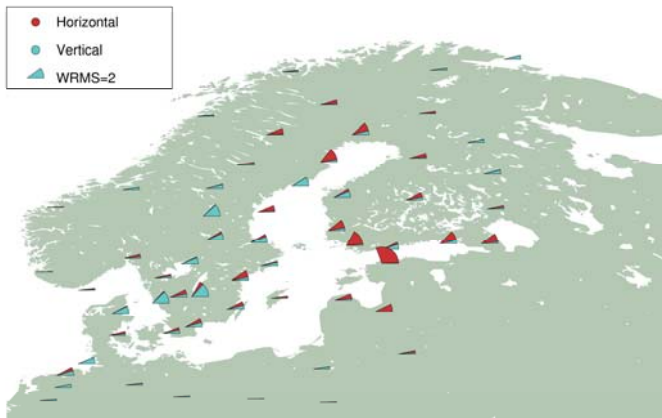
Polyfit vertical



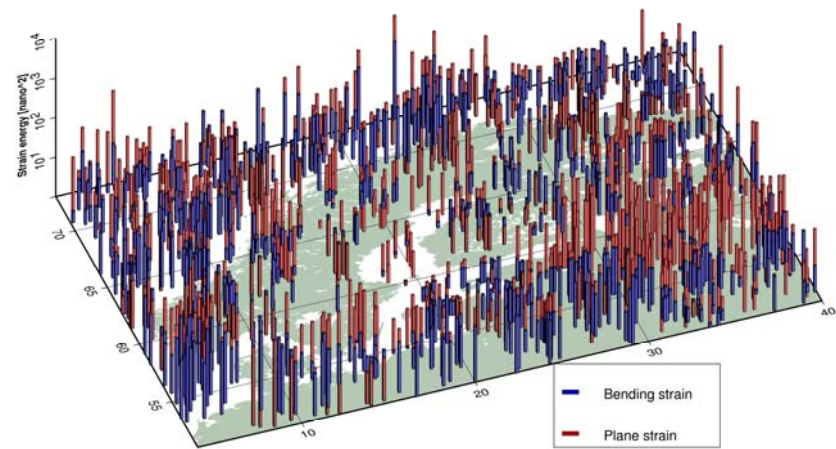
Bending strain rates



Displacement model error



Strain Energy



Thick- (thin-) plate model results of observation fit

Input data	Signal wrms	Resid. Wrms
BIFROST-GPS	15.2	0.78 (0.79)
GIA-model	12.6	0.07 (0.04)
GPS minus GIA-model	4.5	0.90 (0.78)



Conclusions

- Early stage of model development
- Clearly, flexure appears important
- Little signal to be explained after subtraction of GIA-model from BIFROST velocities
- Strain rate field looks smooth except for an anomalous area in Finnmarken
- This area is not well sampled with GPS stations
- In this area also curl is peaking



Deformation energy

Bending: vertical displacement

$$E_b = \frac{D^3}{24} \left\{ [\lambda(1 - \nu)^2 + 2\mu(1 + \nu^2)] (\nabla_h^2 u)^2 - 4 \frac{\mu(1 + \nu)^2}{r^4 \sin^2 \theta} \frac{\partial^2 u}{\partial \theta^2} \frac{\partial^2 u}{\partial \lambda^2} + 2 \frac{\mu}{r^4 \sin^2 \theta} \left(\frac{\partial^2 u}{\partial \theta \partial \lambda} \right)^2 \right\}$$

Horizontal deformation: Mid-plane strain

$$E_p = D \left[\frac{\lambda + 2\mu}{2} (\epsilon_{\lambda\lambda}^2 + \epsilon_{\theta\theta}^2) + \lambda \epsilon_{\theta\theta} \epsilon_{\lambda\lambda} + \mu \epsilon_{\theta\lambda}^2 \right]$$

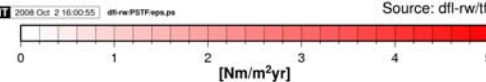
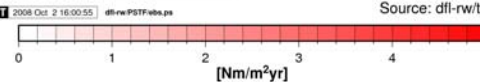
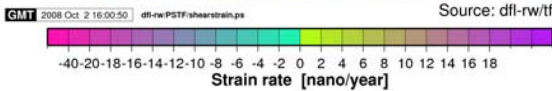
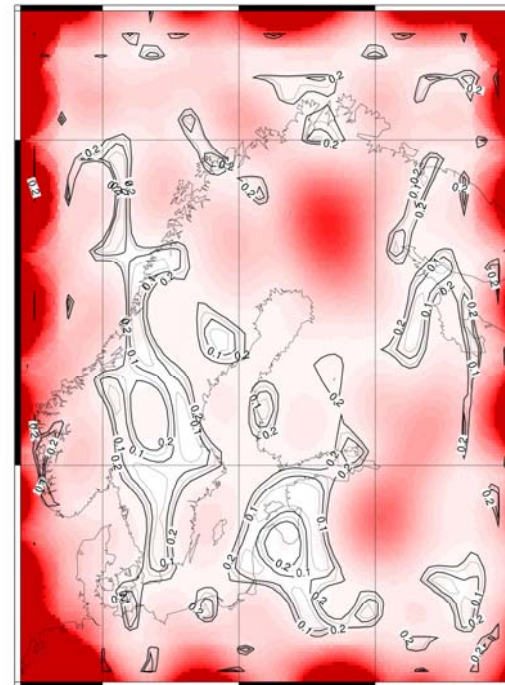
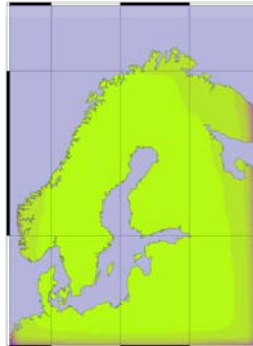
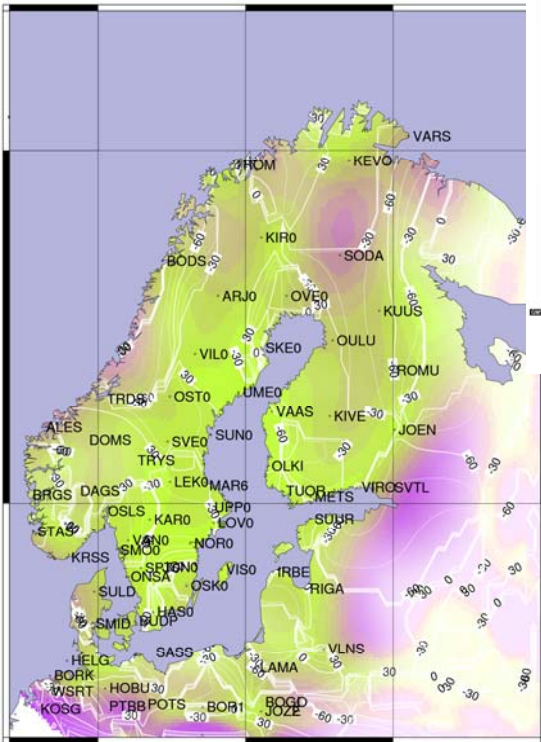


Shear strain

Shear strain

Shear bending strain energy

Shear plane strain energy



Displacement model error

WRMS=2



Strain Energy

